

that we obtain this inversion formula from the input source/receiver configuration and earth model. That is,

$$\beta(\mathbf{x}) = \frac{1}{8\pi^3} \int d^2\xi_I \frac{|h(\mathbf{x}, \boldsymbol{\xi}_I)|}{a_I(\mathbf{x}, \boldsymbol{\xi}_I) |\nabla_x \tau_I(\mathbf{x}, \boldsymbol{\xi}_I)|} \int i\omega_I d\omega_I e^{-i\omega_I \tau_I(\mathbf{x}, \boldsymbol{\xi}_I)} u_I(\boldsymbol{\xi}_I, \omega_I). \quad (35)$$

Our objective is to map the input data set, with its source/receiver configuration and background parameters (macro-model) to an output data set, with its source/receiver configuration and background parameters. To achieve this, the representation (35) is substituted into (34) to obtain the following representation for the mapping of data from any input source/receiver configuration and background model to any output source/receiver configuration and background model:

$$u_O(\boldsymbol{\xi}_O, \omega_O) \sim -\frac{i\omega_O}{8\pi^3} \int i\omega_I d\omega_I d^2\xi_I u_I(\boldsymbol{\xi}_I, \omega_I) \int \frac{a_O(\mathbf{x}, \boldsymbol{\xi}_O)}{a_I(\mathbf{x}, \boldsymbol{\xi}_I)} \frac{|\nabla_x \tau_O(\mathbf{x}, \boldsymbol{\xi}_O)|}{|\nabla_x \tau_I(\mathbf{x}, \boldsymbol{\xi}_I)|} |h(\mathbf{x}, \boldsymbol{\xi}_I)| e^{[i\omega_O \tau_O(\mathbf{x}, \boldsymbol{\xi}_O) - i\omega_I \tau_I(\mathbf{x}, \boldsymbol{\xi}_I)]} d^3x. \quad (36)$$

Note that the input data in the first line here is independent of the earth modeling variables,  $\mathbf{x}$ . Hence, for each choice of input and output earth model and each choice of input and output source/receiver configuration, the integrations over  $\mathbf{x}$  in the second and third lines could be carried out to obtain an operator kernel that is a function of  $\boldsymbol{\xi}_I$ ,  $\omega_I$ ,  $\boldsymbol{\xi}_O$ , and  $\omega_O$ . Indeed, we anticipate carrying out those integrations by analytical methods including asymptotic methods such as multi-dimensional stationary phase. Numerical integration is out of the question. There are  $O(n^3)$  coordinates of integration, with  $O(n^3)$  input variables and  $O(n^3)$  output variables. Clearly, the  $n$ 's are different, but this is still an intractably large set of variables. The processing of this formula would require an integration over  $O(n^3)$  variables to produce a function of  $O(n^6)$  variables. A better choice is the analysis of the volume integral represented by the second and third line to obtain an analytically explicit kernel that depends only on the input and output variables,  $\boldsymbol{\xi}_I$ ,  $\omega_I$ ,  $\boldsymbol{\xi}_O$ , and  $\omega_O$ .

From the derivation and the specific example of 2.5D DMO, it is expected that the result will not transform the reflection coefficient of the input data configuration to the reflection coefficient of the output configuration. However, we do anticipate that the geometrical spreading effects and curvature effects of the input configuration will be transformed to the correct effects for the output configuration. In fact, this has been proven by Tygel, et al, [1998], using a somewhat different (time-domain) approach to the same problem. While their proof does not include all of the cases listed above, it could easily be extended to them. In that sense, we consider their proof as all encompassing for this operator.

There is extensive discussion and interpretation of the structure of this operator in Bleistein [1998a]. Here, we only point out that some implementations have already

been carried out and are described in Bleistein [1998b, c], Sheaffer and Bleistein, [1998]. The 2.5D implementation for mapping from common offset data to common shot data has also been worked out, but not yet presented.

## 2.5D

The specialization of (36) follows from the same type of “thought experiment” as was introduced in the previous section. Except for the subscripts  $O$  and  $I$ , the variables are as described in that section. The main result is the 2.5D platform for datamapping. That result is

$$\begin{aligned}
 u_O(\xi_O, \omega_O) \sim & \frac{\sqrt{|\omega_O|} e^{-i\pi \operatorname{sgn}(\omega_O)/4}}{4\pi^2} \int \sqrt{|\omega_I|} e^{i\pi \operatorname{sgn}(\omega_I)/4} d\omega_I d\xi_I u_I(\xi_I, \omega_I) \\
 & \cdot \int \frac{a_O(\mathbf{x}, \xi_O)}{a_I(\mathbf{x}, \xi_I)} \frac{|\nabla_x \tau_O(\mathbf{x}, \xi_O)|}{|\nabla_x \tau_I(\mathbf{x}, \xi_I)|} \frac{\sqrt{\sigma_{Is} + \sigma_{Ig}}}{\sqrt{\sigma_{Os} + \sigma_{Og}}} \frac{\sqrt{\sigma_{Os} \sigma_{Og}}}{\sqrt{\sigma_{Is} \sigma_{Ig}}} \quad (37) \\
 & \cdot |H(\mathbf{x}, \xi_I)| \cdot e^{[i\omega_O \tau_O(\mathbf{x}, \xi_O) - i\omega_I \tau_I(\mathbf{x}, \xi_I)]} d^2x.
 \end{aligned}$$

This result has been tested for input data represented by the Kirchhoff approximation, in full generality of an unprescribed source/receiver configuration. We find that, asymptotically, the Kirchhoff data in the input source/receiver variables is mapped to a Kirchhoff approximation in the output source/receiver variables, except that the reflection coefficient is evaluated at the specular angle associated with the input variables.

We know that the Kirchhoff approximation contains the “right” geometrical spreading effects, both due the source and receiver Green’s functions and due to the reflector curvature. Therefore, there is no need to check individual implementations separately; the mapping performs as claimed.

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